


2017

Hamiltonian Model for Coupled Surface and Internal Waves in the Presence of Currents

Rossen Ivanov

Technological University Dublin, rossen.ivanov@tudublin.ie

Follow this and additional works at: <https://arrow.tudublin.ie/scschmatart>

 Part of the [Mathematics Commons](#), [Non-linear Dynamics Commons](#), and the [Partial Differential Equations Commons](#)

Recommended Citation

Ivanov, R. (2017) Hamiltonian model for coupled surface and internal waves in the presence of currents. *Nonlinear Analysis: Real World Applications* Volume 34 (2017) Pages 316-334. doi:10.1016/j.nonrwa.2016.09.010

This Article is brought to you for free and open access by the School of Mathematics at ARROW@TU Dublin. It has been accepted for inclusion in Articles by an authorized administrator of ARROW@TU Dublin. For more information, please contact arrow.admin@tudublin.ie, aisling.coyne@tudublin.ie.



This work is licensed under a [Creative Commons Attribution-Noncommercial-Share Alike 4.0 License](#)
Funder: DIT

Hamiltonian model for coupled surface and internal waves in the presence of currents

Rossen Ivanov

*School of Mathematical Sciences, Dublin Institute of Technology,
Kevin Street, Dublin 8, Ireland*

and

*Environmental Sustainability and Health Institute (ESHI), Dublin Institute of Technology
Grangegorman, Dublin 7, Ireland*

Abstract

We examine a two dimensional fluid system consisting of a lower medium bounded underneath by a flatbed and an upper medium with a free surface. The two media are separated by a free common interface. The gravity driven surface and internal water waves (at the common interface between the media) in the presence of a depth-dependent current are studied under certain physical assumptions. Both media are considered incompressible and with prescribed vorticities. Using the Hamiltonian approach the Hamiltonian of the system is constructed in terms of 'wave' variables and the equations of motion are calculated. The resultant equations of motion are then analysed to show that wave-current interaction is influenced only by the current profile in the 'strips' adjacent to the surface and the interface. Small amplitude and long-wave approximations are also presented.

Keywords: Internal waves, Equatorial undercurrent, shear flow, Hamiltonian system, KdV equation

2000 MSC: 35Q35, 37K05, 74J30

1. Introduction

It has been known for many centuries that the ocean contains currents that flow along generally consistent paths. The Spanish galleons transporting gold and silver from Mexico to Spain made use of the Gulf Stream to

help them return home. Since then, scientists have gained much more information on both where currents flow and why. In the oceans currents very often exist with undercurrents. The first undercurrent was discovered in 1951 by Townsend Cromwell who was investigating fishing techniques in the central Pacific Ocean. Undercurrents have since been found under most major currents. The equatorial region in the Pacific is characterised by a thin shallow layer of warm and less dense water over a much deeper layer of cold denser water. The two layers are separated by a sharp *thermocline* (where the temperature gradient has a maximum, it is very close to the *pycnocline*, where the pressure gradient has a maximum) at a depth, depending on the location, but usually at 100 – 200 m beneath the surface. For modelling purposes both layers are assumed homogeneous with a sharp boundary at the thermocline/pycnocline (see [36]).

The Equatorial Undercurrent (EUC) flows in a region that is roughly within 200 - 300 km (below 3° latitude) of the Equator, it is symmetric about the Equator and extends nearly across the whole length (more than 12000 km) of the Pacific Ocean basin [34]. With speeds in excess of 1 m/s, the EUC is one of the fastest permanent currents in the world.

The flow has nearly two-dimensional character, with small meridional variations. While at depths in excess of about 240 m there is, essentially, an abyssal layer of still water, the ocean dynamics near the surface is quite complex. In this region the wave motion typically comprises surface gravity waves with amplitudes of 1-2 m and oscillations with an amplitude of 10-20 m at the thermocline (of mean depth between 50 m and 150 m). These waves interact with the underlying currents. In that case the velocity is (anti-) parallel to the Earth's angular speed ω , so their vector product is zero. This feature distinguishes the dynamics of the equatorial zone from the ocean dynamics at higher latitudes.

The strong stratification confines the wind-driven currents to a shallow near-surface region, less than 200 m deep. In the Atlantic and Pacific, the westward trade winds induce a westward surface flow at speeds of 25-75 cm/s, while a jet-like current – the Equatorial Undercurrent (EUC) – flows below it toward the East (counter to the surface current), attaining speeds of more than 1 m/s at a depth of nearly 100 m. The wind-generated equatorial current in the layer above the thermocline is with a strictly monotonic depth-dependence and exhibits flow-reversal, while beneath the thermocline the current simply decays with increasing depth, being irrelevant in the abyssal region.

While viscous theory is essential in explaining the generation of the equatorial current induced by wind forcing, inviscid theory is adequate for the study of non-turbulent wave-current interactions since the relevant Reynolds numbers are very large (see [43]).

For some general facts concerning the description of waves interacting with currents we refer to the following reviews and monographs [10, 49, 37, 52] and the references therein. The present study draws from previous single medium irrotational [54], [3], [46], [47], [48] and rotational [9], [11], [10], [12], [50], [17], [53], [42] studies as well as from studies of two-media systems such as [1], [2], [22], [21], [18], [19], [16], [15], [4], [5],[6], [7], [20],[27],[28], [29], [41], [44], [45].

The Hamiltonian approach to water waves dynamics has been put forward for the first time by Zakharov [54]. The Hamiltonian formulation describing the two-dimensional nonlinear interaction between coupled surface waves, internal waves, and an underlying current with piecewise constant vorticity, in a two-media fluid overlying a flat bed has been developed in [16], [15]. In the present study we will be following a similar approach, taking into account the shear current structure suggested in [18]. Related results for a flat surface (effectively rigid lid) has been studied in [4] [5],[6], [7], [8].

The model equations will be presented in a canonical Hamiltonian form and then small amplitude and long wave approximations will be derived.

2. Preliminaries

The system under study involves two-dimensional surface and internal gravity water waves and a depth dependent current as per Figure 1.

The medium underneath the internal wave is defined by the domain

$$\Omega(\eta) = \{(x, y) \in \mathbb{R}^2 : -h < y < \eta(x, t)\}.$$

This medium is bounded at the bottom by an impermeable flatbed at a depth $-h$. The medium above the internal wave $y = \eta(x, t)$ is the domain

$$\Omega_1(\eta, \eta_1) = \{(x, y) \in \mathbb{R}^2 : \eta(x, t) < y < h_1 + \eta_1(x, t)\}.$$

This medium is regarded as being bounded on top by a surface wave at a height $y = h_1 + \eta_1(x, t)$ moving around the average level $y = h_1$. Throughout the article the subscript 1 will be used to mean evaluation for the upper medium Ω_1 , and no subscript means evaluation for the lower medium Ω .

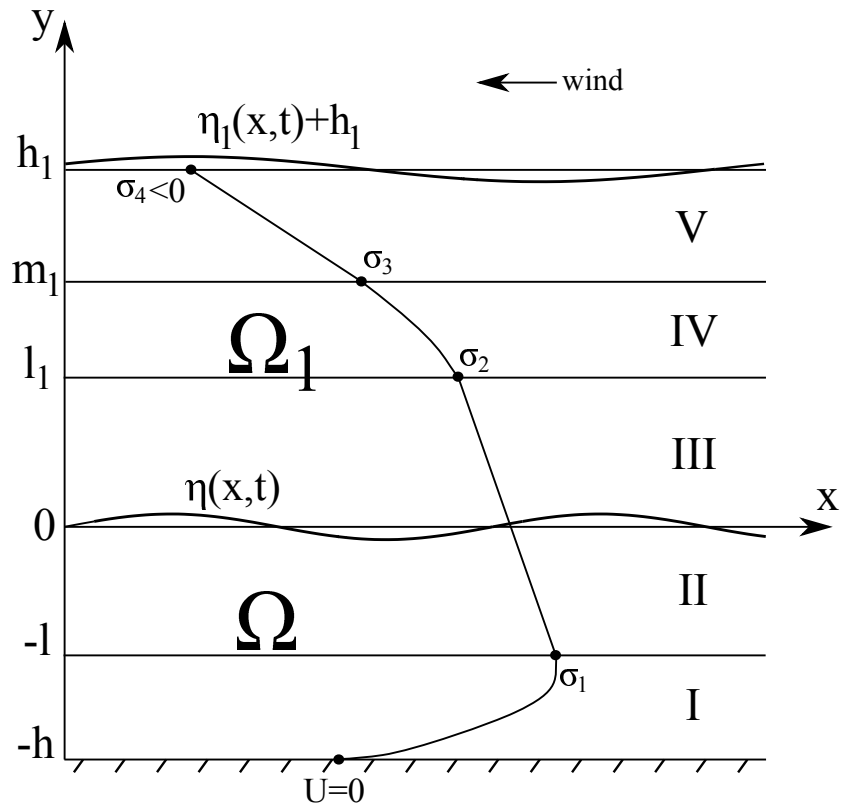


Figure 1: System setup. The current profile in layers I and IV is arbitrary as we are only concerned with layers II, III and V as the internal wave is a free interface between these layers. Continuity of $U(y)$ is assumed in layers I and IV.

Subscript s will be used to denote evaluation at the common interface (thermocline/pycnocline), subscript s_1 - evaluation at the free surface.

The velocity field is denoted by $\mathbf{V}(x, y, z) = (u, v, 0)$ in Ω and $\mathbf{V}_1(x, y, z) = (u_1, v_1, 0)$ in Ω_1 . The function $\eta(x, t)$ describes the deviation of the internal wave from its average level $y = 0$, i.e. $\int_{\mathbb{R}} \eta(x, t) dx = 0$. Similarly, we define the mean of $\eta_1 + h_1$ to be the unperturbed surface $y = h_1$, i.e. $\int_{\mathbb{R}} \eta_1(x, t) dx = 0$.

A depth dependent current $U(y)$ has the following structure:

$$U(y) = \begin{cases} \sigma_4 = \gamma_1 h_1 + \kappa_1 < 0 & y = h_1, \\ \gamma_1 y + \kappa_1, & m_1 \leq y & \text{(layer V)}, \\ \sigma_3 = \gamma_1 m_1 + \kappa_1 & y = m_1, \\ \sigma_2 = \gamma l_1 + \kappa & y = l_1, \\ \gamma y + \kappa, & -l \leq y \leq l_1 & \text{(layers II and III)}, \\ \sigma_1 = -\gamma l + \kappa & y = -l, \\ 0 & y = -h \text{ (flatbed)}, \end{cases} \quad (1)$$

for constants $\sigma_1, \sigma_2, \sigma_3, \sigma_4 < 0$, $\kappa, \kappa_1, l, l_1, \gamma$ and $\gamma_1 < 0$, where κ is the time-independent current velocity at $y = 0$; γ and γ_1 are the non-zero constant vorticity for layers II, III and V, noting that the current is not specified explicitly in layers I and IV, and satisfies only the condition that in the whole fluid body $U(y)$ is a continuous function. This setup of the wave-current system is motivated in [18], see also [6, 7]. The wave motion on the surface is usually confined in the top layer V', defined for depths y such that

$$V' : \quad m_1 \leq y \leq h_1 + \eta_1(x, t),$$

and the internal wave is confined in the strip formed of layers II and III:

$$-l \leq \eta(x, t) \leq l_1.$$

We define also layers II' and III':

$$\begin{aligned} \text{II}' : \quad & -l \leq y \leq \eta(x, t), \\ \text{III}' : \quad & \eta(x, t) \leq y \leq l_1. \end{aligned} \quad (2)$$

We consider a velocity field, which is defined by the wave-related velocity potentials $\varphi(x, y, t)$ for the domain Ω and $\varphi_1(x, y, t)$ for the domain Ω_1 as

follows:

$$\begin{cases} u_1 = \varphi_{1,x} + \gamma_1 y + \kappa_1 & (\text{layer V}'), \\ u_1 = \varphi_{1,x} + \gamma y + \kappa & (\text{layer III}'), \\ v_1 = \varphi_{1,y}, \\ u = \varphi_x + \gamma y + \kappa & (\text{layer II}'), \\ v = \varphi_y. \end{cases} \quad (3)$$

We note that this representation separates the wave and current contributions to the velocity in layers II', III' and V', so that the horizontal velocity field in Ω is nominally separated to a wave and current part, i.e. $u = \varphi_x + U(y)$, and similarly $u_1 = \varphi_{1,x} + U(y)$ for Ω_1 , see also [15, 16].

The respective constant densities ρ and ρ_1 of the lower and upper media satisfy the stability condition for immiscibility

$$\rho > \rho_1. \quad (4)$$

The rotationality of the layers II and III is given by constant vorticity

$$\gamma = \frac{\sigma_2 - \sigma_1}{l + l_1}. \quad (5)$$

The sign of γ is not specified and the case $\gamma = 0$ is of course possible.

We assume that for large $|x|$ the amplitude of η attenuates and hence make the following assumptions

$$\lim_{|x| \rightarrow \infty} \eta(x, t) = 0, \quad \lim_{|x| \rightarrow \infty} \eta_1(x, t) = 0, \quad (6)$$

$$\lim_{|x| \rightarrow \infty} \varphi(x, y, t) = 0, \quad \lim_{|x| \rightarrow \infty} \varphi_1(x, y, t) = 0. \quad (7)$$

Moreover, we assume

$$-l \leq \eta(x, t) < l_1, \quad m_1 \leq h_1 + \eta_1(x, t) \quad (8)$$

for all x and t . This is an assumption that the surface wave motion takes place only in the *strip* of layer V', and the internal wave motion is only in the *strip* of layers II and III.

3. Governing equations

The governing equations for an inviscid flow are the Euler's equations modified with terms, taking into account the Coriolis force.

Denoting with $(\mathbf{u}(x, y, t), \mathbf{v}(x, y, t))$ the velocity field in $\Omega \cup \Omega_1$, the horizontal component of the velocity field is

$$\mathbf{u}(x, y, t) := \begin{cases} u(x, y, t), & \text{in } \Omega(\eta), \\ u_1(x, y, t), & \text{in } \Omega_1(\eta, \eta_1), \end{cases} \quad (9)$$

and the vertical component of the velocity field is

$$\mathbf{v}(x, y, t) := \begin{cases} v(x, y, t), & \text{in } \Omega(\eta), \\ v_1(x, y, t), & \text{in } \Omega_1(\eta, \eta_1). \end{cases} \quad (10)$$

The mass conservation is given by

$$\mathbf{u}_x + \mathbf{v}_y = 0 \text{ in } \Omega \cup \Omega_1. \quad (11)$$

The equation of mass conservation (11) ensures the existence of a stream function

$$\boldsymbol{\psi}(t, x, y) = \begin{cases} \psi(t, x, y) & \text{in } \Omega, \\ \psi_1(t, x, y) & \text{in } \Omega_1, \end{cases}$$

determined up to an additive term that depends only on time, by

$$\begin{cases} u = \psi_y, & v = -\psi_x, & \text{in } \Omega, \\ u_1 = \psi_{1,y}, & v_1 = -\psi_{1,x}, & \text{in } \Omega_1. \end{cases} \quad (12)$$

For convenience we introduce also

$$\boldsymbol{\varphi}(x, y, t) := \begin{cases} \varphi(x, y, t), & \text{in } \Omega(\eta), \\ \varphi_1(x, y, t), & \text{in } \Omega_1(\eta, \eta_1), \end{cases} \quad (13)$$

then

$$\mathbf{u} = \boldsymbol{\varphi}_x + U(y).$$

Considering equatorial motion the following Coriolis forces per unit mass have to be taken into account:

$$\mathbf{F} = 2\omega \nabla \boldsymbol{\psi} = 2\omega(-\mathbf{v}, \mathbf{u})^T \quad (14)$$

with ω being the rotational speed of the Earth. Then the Euler's equations are

$$\begin{cases} \mathbf{u}_t + \mathbf{u}\mathbf{u}_x + \mathbf{v}\mathbf{u}_y + 2\omega\mathbf{v} &= -\frac{1}{\rho^*}P_x, \\ \mathbf{v}_t + \mathbf{u}\mathbf{v}_x + \mathbf{v}\mathbf{v}_y - 2\omega\mathbf{u} &= -\frac{1}{\rho^*}P_y - g, \end{cases} \quad (15)$$

where $P = P(x, y, t)$ denotes the pressure, g is the gravitational acceleration and ρ^* denotes the density of the fluid which, in our case is assumed to be piecewise constant, equal to ρ in $\Omega(\eta)$ and to ρ_1 in $\Omega_1(\eta, \eta_1)$.

Complementing the equations of motion are the boundary conditions, of which

$$P = P_{atm} \text{ on } y = \eta_1(x) + h_1, \quad (16)$$

(with P_{atm} being the constant atmospheric pressure) decouples the motion of the water from that of the air. In addition to (16) we have the kinematic boundary conditions which refer to the flat bed, the interface $y = \eta(x, t)$, the free surface $y = h_1 + \eta_1(x, t)$ and reflect the impermeability of these three surfaces. Thus, they read as

$$v_1 = \eta_{1,t} + u_1\eta_{1,x} \text{ on } y = \eta_1(x, t) + h_1, \quad (17)$$

$$\begin{aligned} v_1 &= \eta_t + u_1\eta_x & \text{on } y &= \eta(x, t), \\ v &= \eta_t + u\eta_x & \text{on } y &= \eta(x, t), \end{aligned} \quad (18)$$

and

$$v = 0 \text{ on } y = -h. \quad (19)$$

From the kinematic boundary conditions (18) one can obtain

$$\psi(t, x, \eta(x, t)) = \psi_1(t, x, \eta(x, t)) = -\int_{-\infty}^x \eta_t(x', t) dx', \quad (20)$$

i.e. ψ is a continuous function across $y = \eta$. This also implies that the normal velocity field components are continuous across the interface $y = \eta(x, t)$.

We introduce for convenience

$$\chi(x, t) \equiv \psi(t, x, \eta(x, t)) = \psi_1(t, x, \eta(x, t)). \quad (21)$$

Similarly

$$\chi_1(x, t) \equiv \psi_1(t, x, \eta_1(x, t) + h_1) = -\int_{-\infty}^x \eta_{1,t}(x', t) dx'. \quad (22)$$

With the velocity decompositions (3) the kinematic boundary conditions (17) and (18) can now be written as

$$\eta_{1,t} = (\varphi_{1,y})_{s_1} - \eta_{1,x}[(\varphi_{1,x})_{s_1} + \gamma_1(h_1 + \eta_1) + \kappa_1] \quad (23)$$

and respectively, as

$$\begin{aligned} \eta_t &= (\varphi_{1,y})_s - \eta_x[(\varphi_{1,x})_s + \gamma\eta + \kappa], \\ \eta_t &= (\varphi_y)_s - \eta_x[(\varphi_x)_s + \gamma\eta + \kappa], \end{aligned} \quad (24)$$

where the subscript s_1 means that we look at traces of the involved functions on the free surface $y = \eta_1(x, t) + h_1$, while the subscript s denotes traces on the interface $y = \eta(x, t)$. The following notation will be used later on in the paper. Namely, we set

$$\begin{aligned} \Phi(x, t) &= \varphi(x, \eta(x, t), t), \\ \Phi_1(x, t) &= \varphi_1(x, \eta(x, t), t), \\ \Phi_2(x, t) &= \varphi_1(x, h_1 + \eta_1(x, t), t). \end{aligned} \quad (25)$$

The Dirichlet-Neumann operator $G(\eta)$ associated to the layer $\Omega(\eta)$ is defined as

$$G(\eta)\Phi := \sqrt{1 + \eta_x^2} \frac{\partial \varphi}{\partial n} \Big|_{y=\eta}, \quad (26)$$

where n denotes the unit outward normal vector to the layer $\Omega(\eta)$ along the interface $y = \eta(x)$. Recall that φ_1 is the solution of the boundary value problem

$$\begin{cases} \Delta \varphi_1 = 0 & \text{in } \Omega(\eta, \eta_1), \\ \varphi_1 = \Phi_1 & \text{on } y = \eta, \\ \varphi_1 = \Phi_2 & \text{on } y = h_1 + \eta_1. \end{cases} \quad (27)$$

The Dirichlet-Neumann operator $G_1(\eta, \eta_1)$ associated to the upper media $\Omega_1(\eta, \eta_1)$ is defined through

$$G_1(\eta, \eta_1)(\Phi_1, \Phi_2) := \begin{pmatrix} -\sqrt{1 + \eta_x^2} \frac{\partial \varphi_1}{\partial n} \Big|_{y=\eta} \\ \sqrt{1 + \eta_{1,x}^2} \frac{\partial \varphi_1}{\partial n_1} \Big|_{y=\eta_1+h_1} \end{pmatrix}, \quad (28)$$

where, we denote by n_1 the unit outward normal vector to $\Omega_1(\eta, \eta_1)$ along the free surface $y = \eta_1(x) + h_1$. Of course, $G_1(\eta, \eta_1)$ is a matrix-operator, for which we choose the notation

$$G_1(\eta, \eta_1) = \begin{pmatrix} G_{11} & G_{12} \\ G_{21} & G_{22} \end{pmatrix}. \quad (29)$$

Euler's equations can be recast by means of the stream function and of the generalized velocity potential in the given layers in the form of Bernoulli conservation laws as follows:

$$\varphi_{1,t} + \frac{1}{2}|\nabla\psi_1|^2 - (\gamma_1 + 2\omega)\psi_1 + \frac{P}{\rho_1} + gy = \tilde{f}_1(t) \quad \text{in layer V'}$$

where $\tilde{f}_1(t)$ is an arbitrary function of t . This is related to the freedom to change φ_1 , if necessary, by a time-dependent factor. Making use of (16) we can absorb the constant P_{atm} in the arbitrary function and obtain on $y = h_1 + \eta_1(x, t)$

$$\rho_1(\varphi_{1,t})_{s_1} + \frac{\rho_1}{2}|\nabla\psi_1|_{s_1}^2 - \rho_1(\gamma_1 + 2\omega)\chi_1 + \rho_1g(h_1 + \eta_1) = f_1(t), \quad (30)$$

for another function

$$f_1(t) = \rho_1\tilde{f}_1(t) - P_{atm}. \quad (31)$$

Similarly, on the interface $y = \eta(x, t)$ the pressure is continuous function and thus it follows that

$$\begin{aligned} & \rho \left[(\varphi_t)_s + \frac{|\nabla\psi|_s^2}{2} - (\gamma + 2\omega)\chi + g\eta \right] \\ &= \rho_1 \left[(\varphi_{1,t})_s + \frac{|\nabla\psi_1|_s^2}{2} - (\gamma + 2\omega)\chi + g\eta \right] + f_2(t), \end{aligned} \quad (32)$$

where $f_2(t)$ is another function of t . The equality (32) can be written as

$$[\rho\varphi_t - \rho_1\varphi_{1,t}]_s + \frac{\rho|\nabla\psi|_s^2}{2} - \frac{\rho_1|\nabla\psi_1|_s^2}{2} - (\rho - \rho_1)(\gamma + 2\omega)\chi + (\rho - \rho_1)g\eta = f_2(t). \quad (33)$$

Introducing the constants

$$\begin{aligned} \Gamma_1 &= \rho_1(\gamma_1 + 2\omega), \\ \Gamma &= (\rho - \rho_1)(\gamma + 2\omega), \end{aligned} \quad (34)$$

the variables (suggested as canonical variables in the Hamiltonian formulation in [1, 2])

$$\begin{aligned} \xi_1(x, t) &= \rho_1\varphi_1(x, \eta_1(x, t) + h_1, t) \equiv \rho_1\Phi_2(x, t), \\ \xi &= \rho\varphi(x, \eta(x, t), t) - \rho_1\varphi_1(x, \eta(x, t), t) \equiv \rho\Phi(x, t) - \rho_1\Phi_1(x, t) \end{aligned} \quad (35)$$

and noting that

$$[\rho\varphi_x - \rho_1\varphi_{1,x}]_s = \xi_x - [\rho\varphi_y - \rho_1\varphi_{1,y}]_s\eta_x, \quad (36)$$

$$[\rho\varphi_t - \rho_1\varphi_{1,t}]_s = \xi_t - [\rho\varphi_y - \rho_1\varphi_{1,y}]_s\eta_t, \quad (37)$$

$$\rho_1(\varphi_{1,t})_{s_1} = \xi_{1,t} - \rho_1(\varphi_{1,y})_{s_1}\eta_{1,t}, \quad (38)$$

$$\rho_1(\varphi_{1,x})_{s_1} = \xi_{1,x} - \rho_1(\varphi_{1,y})_{s_1}\eta_{1,x}, \quad (39)$$

we recast (30) and (33) in terms of the velocity potentials and the current parameters, using the relation between the stream function and the velocity potentials that follows from (12) and (3):

$$\begin{aligned} \xi_{1,t} + \frac{\rho_1}{2}|\nabla\varphi_1|_{s_1}^2 - \rho_1(\varphi_{1,y})_{s_1}[\varphi_{1,y} - \varphi_{1,x}\eta_{1,x}]_{s_1} + \rho_1g(h_1 + \eta_1) \\ + (\gamma_1(\eta_1 + h_1) + \kappa_1)\xi_{1,x} + \frac{\rho_1}{2}(\gamma_1(\eta_1 + h_1) + \kappa_1)^2 - \Gamma_1\chi_1 = f_1(t), \end{aligned} \quad (40)$$

$$\begin{aligned} \xi_t - [\rho\varphi_y - \rho_1\varphi_{1,y}]_s[\varphi_y - \varphi_x\eta_x]_s + \frac{\rho|\nabla\varphi|_s^2}{2} - \frac{\rho_1|\nabla\varphi_1|_s^2}{2} + (\rho - \rho_1)g\eta \\ + (\gamma\eta + \kappa)\xi_x + \frac{\rho - \rho_1}{2}(\gamma\eta + \kappa)^2 - \Gamma\chi = f_2(t). \end{aligned} \quad (41)$$

The balance of all quantities when $x \rightarrow \pm\infty$ gives

$$f_1 = \rho_1 \left(gh_1 + \frac{1}{2}(\gamma_1 h_1 + \kappa_1)^2 \right), \quad f_2 = \frac{1}{2}(\rho - \rho_1)\kappa^2. \quad (42)$$

Recall that in the absence of current and Coriolis force the system is Hamiltonian and can be represented in the form

$$\begin{aligned} \xi_{1,t} = -\delta_{\eta_1} H_0, \quad \xi_t = -\delta_{\eta} H_0, \\ \eta_{1,t} = \delta_{\xi_1} H_0, \quad \eta_t = \delta_{\xi} H_0, \end{aligned} \quad (43)$$

where $H_0(\eta, \eta_1, \xi, \xi_1)$ is the corresponding Hamiltonian. H_0 is evaluated in terms of the canonical variables, see [22, 21] for details:

$$\begin{aligned} H_0(\eta, \eta_1, \xi, \xi_1) = \frac{1}{2} \int_{\mathbb{R}} \begin{pmatrix} \xi \\ \xi_1 \end{pmatrix}^t \begin{pmatrix} G_{11}B^{-1}G(\eta) & -G(\eta)B^{-1}G_{12} \\ -G_{21}B^{-1}G(\eta) & -\frac{\rho}{\rho_1}G_{21}B^{-1}G_{12} + \frac{1}{\rho_1}G_{22} \end{pmatrix} \begin{pmatrix} \xi \\ \xi_1 \end{pmatrix} dx \\ + \frac{1}{2} \int_{\mathbb{R}} (g(\rho - \rho_1)\eta^2 + g\rho_1\eta_1^2 + 2g\rho_1h_1\eta_1) dx, \end{aligned}$$

where the operator $B \equiv \rho_1 G(\eta) + \rho G_{11}$.

Hence we can represent all terms not related to vorticity as variational derivatives of H_0 :

$$\begin{aligned} \xi_{1,t} + \delta_{\eta_1} H_0 + (\gamma_1(\eta_1 + h_1) + \kappa_1)\xi_{1,x} + \frac{\rho_1}{2} (\gamma_1(\eta_1 + h_1) + \kappa_1)^2 - \Gamma_1 \chi_1 &= f_1(t), \\ \xi_t + \delta_\eta H_0 + (\gamma\eta + \kappa)\xi_x + \frac{\rho - \rho_1}{2} (\gamma\eta + \kappa)^2 - \Gamma\chi &= f_2(t). \end{aligned} \quad (44)$$

In addition, (23), (24) can be written as

$$\begin{aligned} \eta_{1,t} &= \delta_{\xi_1} H_0 - [\gamma_1(h_1 + \eta_1) + \kappa_1]\eta_{1,x}, \\ \eta_t &= \delta_\xi H_0 - (\gamma\eta + \kappa)\eta_x. \end{aligned} \quad (45)$$

Now we are in a position to write the equations (44), (45) in the form

$$\begin{aligned} \xi_{1,t} &= -\delta_{\eta_1} H + \Gamma_1 \chi_1, & \xi_t &= -\delta_\eta H + \Gamma\chi, \\ \eta_{1,t} &= \delta_{\xi_1} H, & \eta_t &= \delta_\xi H, \end{aligned} \quad (46)$$

where

$$\begin{aligned} H(\eta, \eta_1, \xi, \xi_1) &= H_0 - \kappa \int_{\mathbb{R}} \xi \eta_x dx - \int_{\mathbb{R}} \gamma \eta \eta_x \xi dx + \frac{\rho - \rho_1}{6\gamma} \int_{\mathbb{R}} (\gamma\eta + \kappa)^3 dx \\ &\quad - \kappa_1 \int_{\mathbb{R}} \xi_1 \eta_{1,x} dx - \int_{\mathbb{R}} \gamma_1(\eta_1 + h_1) \eta_{1,x} \xi_1 dx + \frac{\rho_1}{6\gamma_1} \int_{\mathbb{R}} (\gamma_1(\eta_1 + h_1) + \kappa_1)^3 dx \\ &\quad - f_1(t) \int_{\mathbb{R}} \eta_1 dx - f_2(t) \int_{\mathbb{R}} \eta dx - H_{00}. \end{aligned} \quad (47)$$

$$H_{00} = \frac{\rho - \rho_1}{6\gamma} \int_{\mathbb{R}} \kappa^3 dx + \frac{\rho_1}{6\gamma_1} \int_{\mathbb{R}} (\gamma_1 h_1 + \kappa_1)^3 dx$$

is an integral with a constant Hamiltonian density and zero variational derivatives which keeps the overall Hamiltonian density decaying to zero at $x \rightarrow \pm\infty$. The *ghost* terms with $f_{1,2}$ are un-physical (i.e. their values do not affect

the measurable physical quantities such as velocities and elevations), since by definition

$$\int_{\mathbb{R}} \eta dx = 0, \quad \int_{\mathbb{R}} \eta_1 dx = 0.$$

Nevertheless their variational derivatives produce the $f_{1,2}$ terms in (44) which are also un-physical and can in principle be absorbed in the definition of the potentials ξ , ξ_1 , in our case, in such a way that ξ , ξ_1 tend to zero when $x \rightarrow \pm\infty$. From (42) it follows that H does not contain terms, linear in the field variables. The lowest order terms are quadratic in the field variables and they produce the linearised equations.

The expression for the Hamiltonian (47) can be obtained alternatively by evaluation of the total energy of the fluid (up to a constant Hamiltonian density):

$$H = \int \int_{\Omega \cup \Omega_1} \rho^* \left\{ \frac{\mathbf{u}^2 + \mathbf{v}^2}{2} + gy \right\} dy dx.$$

Taking into account the stratification of the fluid the above expression can be rewritten as

$$\begin{aligned} H = & \frac{1}{2} \int_{\mathbb{R}} \int_{-h}^{\eta(x,t)} \rho (u^2 + v^2) dy dx' + \frac{1}{2} \int_{\mathbb{R}} \int_{\eta(x,t)}^{h_1 + \eta_1(x,t)} \rho_1 (u_1^2 + v_1^2) dy dx' \\ & + \int_{\mathbb{R}} \int_{-h}^{\eta(x,t)} g \rho y dy dx' + \int_{\mathbb{R}} \int_{\eta(x,t)}^{h_1 + \eta_1(x,t)} g \rho_1 y dy dx'. \end{aligned} \quad (48)$$

The computations follow the routine from [16, 6, 7]. We mention only that due to the two-dimensional character of the dynamics, the final expression (47) depends only on the variables on the surface and on the interface. It is an important feature, that there is no contribution from layers I and IV where the wave motion on s and s_1 does not take place [6, 7]. The terms related to the internal wave are the same as in [7] where flat surface approximation is considered.

4. Hamiltonian dynamics

The ideas for the Hamiltonian formulation of water waves coupled to a flow with a constant vorticity originate from [17] followed by [53]. For internal waves with vorticity the problem is studied in [16, 15, 4, 5].

The evolution equations (46) can be written in a canonical Hamiltonian form in terms of the variables

$$\begin{aligned} z &= \xi + \frac{\Gamma}{2} \int_{-\infty}^x \eta(x', t) dx', \\ z_1 &= \xi_1 + \frac{\Gamma_1}{2} \int_{-\infty}^x \eta_1(x', t) dx'. \end{aligned} \quad (49)$$

The system described by the phase space variables η, η_1, z, z_1 is Hamiltonian. More precisely,

$$\begin{aligned} \frac{\delta H}{\delta \eta} &= -z_t, & \frac{\delta H}{\delta z} &= \eta_t, \\ \frac{\delta H}{\delta \eta_1} &= -z_{1,t}, & \frac{\delta H}{\delta z_1} &= \eta_{1,t}. \end{aligned}$$

The proof follows the lines of the one in [16]. Since

$$\begin{aligned} \chi(x, t) &= - \int_{-\infty}^x \eta_t(x', t) dx' = \int_{-\infty}^x \frac{\delta H}{\delta \eta(x', t)} dx', \\ \chi_1(x, t) &= - \int_{-\infty}^x \eta_{1,t}(x', t) dx' = \int_{-\infty}^x \frac{\delta H}{\delta \eta_1(x', t)} dx', \end{aligned} \quad (50)$$

the system of equations (46) can be written as

$$\begin{aligned} \dot{\xi}_k &= - \frac{\delta H}{\delta \eta_k} - \Gamma_k \int_{-\infty}^x \frac{\delta H}{\delta \xi_k(x')} dx', \\ \dot{\eta}_k &= \frac{\delta H}{\delta \xi_k}, \end{aligned} \quad (51)$$

where for convenience $k = 0, 1$, $\eta_0 \equiv \eta$, $\xi_0 \equiv \xi$, $\chi_0 \equiv \chi$ and $\Gamma_0 \equiv \Gamma$. This is an equivalent Hamiltonian form, in terms of the original variables and with respect to the Poisson bracket (PB)

$$\begin{aligned} \{A, B\} &= \sum_{k=0}^1 \int_{\mathbb{R}} \left(\frac{\delta A}{\delta \eta_k(x)} \frac{\delta B}{\delta \xi_k(x)} - \frac{\delta A}{\delta \xi_k(x)} \frac{\delta B}{\delta \eta_k(x)} \right) dx \\ &\quad - \Gamma_k \int_{\mathbb{R}} \left(\frac{\delta A}{\delta \xi_k(x)} \int_{-\infty}^x \frac{\delta B}{\delta \xi_k(x')} dx' \right) dx \end{aligned} \quad (52)$$

i.e.

$$\begin{aligned}\dot{\xi}_k &= \{\xi_k, H\} & k = 0, 1, \\ \dot{\eta}_k &= \{\eta_k, H\} & k = 0, 1.\end{aligned}\tag{53}$$

Here we have to specify the spaces where the functionals A, B may belong. The antisymmetry of the PB requires for example

$$\int_{\mathbb{R}} \left(\frac{\delta A}{\delta \xi_k(x)} \int_{-\infty}^x \frac{\delta B}{\delta \xi_k(x')} dx' \right) dx = - \int_{\mathbb{R}} \left(\frac{\delta B}{\delta \xi_k(x)} \int_{-\infty}^x \frac{\delta A}{\delta \xi_k(x')} dx' \right) dx \tag{54}$$

which is possible (e.g. due to integration by parts) if and only if

$$\int_{\mathbb{R}} \frac{\delta A}{\delta \xi_k(x)} dx \int_{\mathbb{R}} \frac{\delta B}{\delta \xi_k(x')} dx' = 0.\tag{55}$$

Thus, at least one of the functionals in the PB should satisfy

$$\int_{\mathbb{R}} \frac{\delta A}{\delta \xi_k(x)} dx = 0.\tag{56}$$

Since in (53) one of the functionals is always H and

$$\int_{\mathbb{R}} \frac{\delta H}{\delta \xi_k(x)} dx = \int_{\mathbb{R}} \partial_t \eta_k(x, t) = \lim_{x \rightarrow \infty} \chi_k(x, t) = 0.$$

5. The pressure in the body of the fluid

The pressure in the body of the fluid can be evaluated from the functions φ and ψ . They can be recovered from ξ and $\xi_1(x, t)$. In addition, there is an interdependency between φ and ψ . since

$$\varphi + i \left(\psi - \int_{-h}^y U(y') dy' \right)$$

is an analytic function of the variable $z = x + iy$ in the domain $\Omega_1 \cup \Omega$. Thus, the corresponding analytic functions in Ω and Ω_1 can be recovered from their values at the boundaries of Ω and Ω_1 , i.e. from $\Phi(x, t)$, $\Phi_1(x, t)$ and $\Phi_2(x, t)$.

From the definition of the Dirichlet-Neumann operators and (24) we have that

$$G_{11}\Phi_1 + G_{12}\Phi_2 = \varphi_{1,x}\eta_x - \varphi_{1,y} = -(\eta_t + (\gamma\eta + \kappa)\eta_x),$$

and

$$G(\eta)\Phi = -\varphi_x\eta_x + \varphi_y = \eta_t + (\gamma\eta + \kappa)\eta_x.$$

Adding up the previous two relations we obtain

$$G_{11}\Phi_1 + G_{12}\Phi_2 + G\Phi = 0. \quad (57)$$

Using (57), and $B := \rho_1 G(\eta) + \rho G_{11}$ and recalling that

$$\xi = \rho\Phi - \rho_1\Phi_1, \quad \xi_1 = \rho_1\Phi_2$$

we can write Φ, Φ_1, Φ_2 in terms of the Hamiltonian variables ξ and ξ_1 as follows

$$\Phi = B^{-1}(G_{11}\xi - G_{12}\xi_1), \quad (58)$$

$$\Phi_1 = B^{-1}\left(-G(\eta)\xi - \frac{\rho}{\rho_1}G_{12}\xi_1\right), \quad (59)$$

$$\Phi_2 = \frac{1}{\rho_1}\xi_1. \quad (60)$$

In the layers with a fixed constant vorticity γ we have Bernoulli conservation laws

$$\varphi_t + \frac{1}{2}|\nabla\psi|^2 + \frac{P}{\rho^*} - (\gamma + 2\omega)\psi + gy = \tilde{\mathbf{f}}(t). \quad (61)$$

From (31) and (42) for the layer V' we have

$$\tilde{\mathbf{f}} = \tilde{f}_1 = \frac{f_1 + P_{atm}}{\rho_1} = \frac{P_{atm}}{\rho_1} + gh_1 + \frac{1}{2}(\gamma_1 h_1 + \kappa_1)^2 \quad (62)$$

Another possible derivation of (62) is the following one. We can evaluate \tilde{f}_1 using the asymptotic values in (61) at $x \rightarrow \pm\infty$ where no wave motion takes place ($y = h_1$) and thus $\psi_{1,x} \rightarrow 0$ and $\psi_{1,y} \rightarrow \gamma_1 h_1 + \kappa_1$ are given by the velocity of the current. Moreover, ψ_1 on the surface is zero when $x \rightarrow \pm\infty$ due to (22). Hence

$$P(x, y, t) = P_{atm} - \rho_1 \left(\varphi_{1t} + \frac{1}{2}|\nabla\psi_1|^2 - (\gamma_1 + 2\omega)\psi_1 + g(y - h_1) \right) + \frac{\rho_1}{2}(\gamma_1 h_1 + \kappa_1)^2 \quad \text{for } m_1 \leq y \leq h_1 + \eta_1(x, t). \quad (63)$$

We can apply the same approach in deriving the pressure in layers III' and II'. At $y = 0$ and $x \rightarrow \pm\infty$ due to the lack of a wave motion the pressure is $P_0 = P_{atm} + \rho_1 g h_1$. For the layer III' we have

$$P(x, y, t) = P_{atm} - \rho_1 \left(\varphi_{1t} + \frac{1}{2} |\nabla \psi_1|^2 - (\gamma + 2\omega)\psi_1 + g(y - h_1) \right) + \frac{\rho_1}{2} \kappa^2$$

for $\eta(x, t) \leq y \leq l_1$,

(64)

For the layer II',

$$P(x, y, t) = P_{atm} + \rho_1 g h_1 - \rho \left(\varphi_t + \frac{1}{2} |\nabla \psi|^2 - (\gamma + 2\omega)\psi + g y \right) + \frac{\rho}{2} \kappa^2$$

for $-l \leq y \leq \eta(x, t)$.

(65)

6. Scales

Let us introduce non-dimensional variables (without bars) related to the dimensional (barred) as follows:

$$\begin{aligned} \bar{t} &= \frac{h_1}{\sqrt{gh_1}} t, & \bar{x} &= h_1 x, & \bar{y} &= h_1 y, & \bar{\eta} &= a \eta, & \bar{\eta}_1 &= a \eta_1, \\ \bar{u} &= \sqrt{gh_1} u, & \bar{u}_1 &= \sqrt{gh_1} u_1, & \bar{v} &= \sqrt{gh_1} v, & \bar{\kappa} &= \sqrt{gh_1} \kappa, \\ \bar{\kappa}_1 &= \sqrt{gh_1} \kappa_1, & \bar{\gamma} &= \frac{\sqrt{gh_1}}{h_1} \gamma, & \bar{\gamma}_1 &= \frac{\sqrt{gh_1}}{h_1} \gamma_1, & \varepsilon &= \frac{a}{h_1}. \end{aligned}$$
(66)

The constant a represents the average amplitude of the waves under consideration, ε is a small parameter which will be used to separate the order of the terms in the model.

From $\bar{v} = \bar{\eta}_{\bar{t}} + \bar{u} \bar{\eta}_{\bar{x}}$ it follows

$$v = \varepsilon(\eta_t + u \eta_x),$$
(67)

therefore, if $\mathcal{O}(\eta_t) = 1$ then $\mathcal{O}(v) = \varepsilon$ and thus the dimensional expression with $\mathcal{O}(v) = 1$ (and similar for v_1) should be

$$\bar{v} = \varepsilon \sqrt{gh_1} v, \quad \bar{v}_1 = \varepsilon \sqrt{gh_1} v_1.$$
(68)

Since v is a y -derivative of the velocity potential, and with the adopted definitions $\bar{\varphi} = \varepsilon h_1 \sqrt{gh_1} \varphi$ etc. and thus

$$\bar{\xi} = \varepsilon \rho h_1 \sqrt{gh_1} \xi, \quad \bar{\xi}_1 = \varepsilon \rho h_1 \sqrt{gh_1} \xi_1. \quad (69)$$

The scales for u , u_1 etc. do not change - their dominant parts are the vorticity and current components of order 1; only the 'wave' component (which is x derivative of φ) is of order ε . The Dirichlet-Neumann operators have the following structure (e.g. any of the introduced operators G , G_{ij}):

$$\bar{G} = \bar{G}^{(0)} + \bar{G}^{(1)} + \bar{G}^{(2)} + \dots \quad (70)$$

where $\bar{G}^{(n)} \sim \bar{\eta}^n \partial_x^{n+1}$ i.e. $\bar{G}^{(n)} = \frac{\varepsilon^n}{h_1} G^{(n)}$:

$$\bar{G} = \frac{1}{h_1} (G^{(0)} + \varepsilon G^{(1)} + \varepsilon^2 G^{(2)} + \dots). \quad (71)$$

With this scaling and ignoring the linear terms, whose average is 0, the Hamiltonian can be expanded as

$$\bar{H} = \rho g h_1^3 (\varepsilon^2 H^{(2)} + \varepsilon^3 H^{(3)} + \dots) \quad (72)$$

where

$$\begin{aligned} H^{(2)} = & \frac{1}{2} \int_{\mathbb{R}} \begin{pmatrix} \xi \\ \xi_1 \end{pmatrix}^t \begin{pmatrix} G_{11} B^{-1} G(\eta) & -G(\eta) B^{-1} G_{12} \\ -G_{21} B^{-1} G(\eta) & -\frac{\rho}{\rho_1} G_{21} B^{-1} G_{12} + \frac{1}{\rho_1} G_{22} \end{pmatrix}^{(0)} \begin{pmatrix} \xi \\ \xi_1 \end{pmatrix} dx \\ & + \frac{1}{2} (\rho - \rho_1) (g + \gamma \kappa) \int_{\mathbb{R}} \eta^2 dx + \frac{1}{2} \rho_1 (g + \gamma_1^2 h_1 + \gamma_1 \kappa_1) \int_{\mathbb{R}} \eta_1^2 dx \\ & - \kappa \int_{\mathbb{R}} \xi \eta_x dx - (\kappa_1 + \gamma_1 h_1) \int_{\mathbb{R}} \xi_1 \eta_{1,x} dx \end{aligned} \quad (73)$$

where the leading order, $\mathcal{O}(1)$ expression for the operators is (does not depend on η , η_1)

$$\begin{pmatrix} G_{11} & G_{12} \\ G_{21} & G_{22} \end{pmatrix}^{(0)} = \begin{pmatrix} D \coth(h_1 D) & -D \operatorname{csch}(h_1 D) \\ -D \operatorname{csch}(h_1 D) & D \coth(h_1 D) \end{pmatrix}, \quad (74)$$

where $D = -i\partial_x$. The quadratic part produces the linearised equations. Similarly

$$\begin{aligned}
H^{(3)} &= \frac{1}{2} \int_{\mathbb{R}} \begin{pmatrix} \xi \\ \xi_1 \end{pmatrix}^t \begin{pmatrix} G_{11}B^{-1}G(\eta) & -G(\eta)B^{-1}G_{12} \\ -G_{21}B^{-1}G(\eta) & -\frac{\rho}{\rho_1}G_{21}B^{-1}G_{12} + \frac{1}{\rho_1}G_{22} \end{pmatrix}^{(1)} \begin{pmatrix} \xi \\ \xi_1 \end{pmatrix} dx \\
&\quad - \gamma \int_{\mathbb{R}} \xi \eta \eta_x dx - \gamma_1 \int_{\mathbb{R}} \xi_1 \eta_1 \eta_{1,x} dx + \frac{1}{2} \int_{\mathbb{R}} \left((\rho - \rho_1) \gamma^2 \frac{\eta^3}{3} + \rho_1 \gamma_1^2 \frac{\eta_1^3}{3} \right) dx.
\end{aligned} \tag{75}$$

The order ε terms $G^{(1)}$ of the operators are given in [22, 21].

7. Linearised equations

The Hamiltonian equations with a Hamiltonian $H^{(2)}$ are the linearised equations:

$$\begin{aligned}
\xi_t &= -\kappa \xi_x - (\rho - \rho_1)(g + \gamma \kappa) \eta - \Gamma \partial^{-1} \eta_t, \\
\eta_t &= -\kappa \eta_x + \frac{D \tanh(hD) \coth(h_1 D)}{\rho \coth(h_1 D) + \rho_1 \tanh(hD)} \xi + \frac{D \tanh(hD) \operatorname{csch}(h_1 D)}{\rho \coth(h_1 D) + \rho_1 \tanh(hD)} \xi_1, \\
\xi_{1,t} &= -(\gamma_1 h_1 + \kappa_1) \xi_{1,x} - \rho_1 (g + \gamma_1^2 h_1 + \gamma_1 \kappa_1) \eta_1 - \Gamma_1 \partial^{-1} \eta_{1,t}, \\
\eta_{1,t} &= -(\gamma_1 h_1 + \kappa_1) \eta_{1,x} + \frac{D \tanh(hD) \operatorname{csch}(h_1 D)}{\rho \coth(h_1 D) + \rho_1 \tanh(hD)} \xi \\
&\quad + \frac{D (\tanh(hD) \coth(h_1 D) + \frac{\rho}{\rho_1})}{\rho \coth(h_1 D) + \rho_1 \tanh(hD)} \xi_1,
\end{aligned} \tag{76}$$

where ∂^{-1} is the inverse of ∂_x . We can change the coordinates via a linear transformation according to $\partial_T = \partial_t + \kappa \partial_x$:

$$\begin{aligned}
\xi_T &= -(\rho - \rho_1)(g - 2\omega \kappa) \eta - (\rho - \rho_1)(\gamma + 2\omega) \partial^{-1} \eta_T, \\
\eta_T &= \frac{D \tanh(hD) \coth(h_1 D)}{\rho \coth(h_1 D) + \rho_1 \tanh(hD)} \xi + \frac{D \tanh(hD) \operatorname{csch}(h_1 D)}{\rho \coth(h_1 D) + \rho_1 \tanh(hD)} \xi_1, \\
\xi_{1,T} &= -(\gamma_1 h_1 + \kappa_1 - \kappa) \xi_{1,x} - \rho_1 [g + \gamma_1^2 h_1 + \gamma_1 (\kappa_1 - \kappa) - 2\omega \kappa] \eta_1 \\
&\quad - \rho_1 (\gamma_1 + 2\omega) \partial^{-1} \eta_{1,T}, \\
\eta_{1,T} &= -(\gamma_1 h_1 + \kappa_1 - \kappa) \eta_{1,x} + \frac{D \tanh(hD) \operatorname{csch}(h_1 D)}{\rho \coth(h_1 D) + \rho_1 \tanh(hD)} \xi \\
&\quad + \frac{D (\tanh(hD) \coth(h_1 D) + \frac{\rho}{\rho_1})}{\rho \coth(h_1 D) + \rho_1 \tanh(hD)} \xi_1.
\end{aligned} \tag{77}$$

Note that the equations contain Coriolis terms dependent on ω . Usually κ is of magnitude several m/s, $2\omega = 1.46 \times 10^{-4} \text{ s}^{-1}$ and hence $2\omega\kappa \ll g$:

$$\begin{aligned}
\xi_T &= -(\rho - \rho_1)g\eta - (\rho - \rho_1)(\gamma + 2\omega)\partial^{-1}\eta_T, \\
\eta_T &= \frac{D \tanh(hD) \coth(h_1D)}{\rho \coth(h_1D) + \rho_1 \tanh(hD)}\xi + \frac{D \tanh(hD) \text{csch}(h_1D)}{\rho \coth(h_1D) + \rho_1 \tanh(hD)}\xi_1, \\
\xi_{1,T} &= -a\xi_{1,x} - \rho_1(g + a\gamma_1)\eta_1 - \rho_1(\gamma_1 + 2\omega)\partial^{-1}\eta_{1,T}, \\
\eta_{1,T} &= -a\eta_{1,x} + \frac{D \tanh(hD) \text{csch}(h_1D)}{\rho \coth(h_1D) + \rho_1 \tanh(hD)}\xi \\
&\quad + \frac{D(\tanh(hD) \coth(h_1D) + \frac{\rho}{\rho_1})}{\rho \coth(h_1D) + \rho_1 \tanh(hD)}\xi_1,
\end{aligned} \tag{78}$$

where we introduced the notation

$$a = \gamma_1 h_1 + \kappa_1 - \kappa.$$

Next we search for solutions, proportional to

$$e^{i(kx - \Omega_0(k)T)}, \tag{79}$$

where $\Omega_0(k)$ is the dispersion law for the wave, the wave speed is

$$c_0(k) = \frac{\Omega_0(k)}{k}.$$

Such a solution would be an eigenfunction for D with a corresponding eigenvalue k . From the system (78) one can express ξ , ξ_1 as

$$\begin{aligned}
\xi &= \frac{i(\rho - \rho_1)}{k} \left(\gamma + 2\omega - \frac{g}{c_0(k)} \right) \eta, \\
\xi_1 &= \frac{i\rho_1}{k} \left(\gamma_1 + 2\omega + \frac{2a\omega - g}{c_0(k) - a} \right) \eta_1.
\end{aligned} \tag{80}$$

The remaining two equations lead to a linear homogeneous system for η and η_1 . We introduce the notations

$$\begin{aligned}
\mu(k) &= \frac{\rho \tanh(hk) \coth(h_1 k)}{\rho \coth(h_1 k) + \rho_1 \tanh(hk)}, \\
f(k) &= \frac{\rho \tanh(hk) \operatorname{csch}(h_1 k)}{\rho \coth(h_1 k) + \rho_1 \tanh(hk)}, \\
\theta(k) &= \frac{\rho \left(\tanh(hk) \coth(h_1 k) + \frac{\rho}{\rho_1} \right)}{\rho \coth(h_1 k) + \rho_1 \tanh(hk)}.
\end{aligned} \tag{81}$$

The second equation of (78) gives

$$\left[c_0 + \frac{(\rho - \rho_1)\mu(k)}{\rho k} \left(\gamma + 2\omega - \frac{g}{c_0} \right) \right] \eta + \frac{\rho_1 f(k)}{\rho k} \left(\gamma_1 + 2\omega + \frac{2a\omega - g}{c_0 - a} \right) \eta_1 = 0. \tag{82}$$

Then the last equation of (78) leads to

$$\frac{(\rho - \rho_1)f(k)}{\rho k} \left(\gamma + 2\omega - \frac{g}{c_0} \right) \eta + \left[c_0 - a + \frac{\rho_1 \theta(k)}{\rho k} \left(\gamma_1 + 2\omega + \frac{2a\omega - g}{c_0 - a} \right) \right] \eta_1 = 0. \tag{83}$$

The compatibility of the two equations gives a 4-th order equation for $c_0(k)$:

$$\begin{aligned}
& \left[c_0 + \frac{\rho - \rho_1}{\rho} \frac{\mu(k)}{k} \left(\gamma + 2\omega - \frac{g}{c_0} \right) \right] \left[c_0 - a + \frac{\rho_1}{\rho} \frac{\theta(k)}{k} \left(\gamma_1 + 2\omega + \frac{2a\omega - g}{c_0 - a} \right) \right] \\
&= \frac{\rho_1(\rho - \rho_1)}{\rho^2} \frac{f^2(k)}{k^2} \left(\gamma + 2\omega - \frac{g}{c_0} \right) \left(\gamma_1 + 2\omega + \frac{2a\omega - g}{c_0 - a} \right).
\end{aligned} \tag{84}$$

The last formula generalises the irrotational one from [21]. In addition, from (83), (82) one can determine if η and η_1 have the same or an opposite polarity for each possible propagation speed c_0 . The velocity $c_0(k)$ is relative to an observer moving together with the flow at $y = 0$, i.e. with velocity κ . The wave speed for a stationary observer therefore is

$$c(k) = c_0(k) + \kappa.$$

8. Long waves approximation

In the long waves approximation the physical scales are measured by the dimensionless parameter like $\delta = \frac{h_1}{L}$. We will study the equations under the additional approximation that the wavelengths L are much bigger than h and h_1 . Since

$$\bar{L} = h_1 L \Rightarrow \frac{1}{L} = \frac{h_1}{\bar{L}} = \delta.$$

Thus the wave number $k = 2\pi/L = 2\pi\delta$ and $\mathcal{O}(k) = \delta$. We further assume that $\delta^2 = \mathcal{O}(\varepsilon)$. Recall that the operator D has an eigenvalue k , thus we shall keep in mind that $\mathcal{O}(D) = \delta$. Moreover the x -derivative of the velocity potentials do not get an extra factor of δ since $\mathcal{O}(\bar{v}) = \varepsilon$ remains unchanged and similarly v_1 . In other words the 'wave' component of u let's call it $\tilde{u} = \varphi_x$ is of order ε and similarly $\tilde{u}_1 = \varphi_{1,x}$ is of order ε . Despite the assumption $\delta^2 = \mathcal{O}(\varepsilon)$ we will keep both scales δ and ε in order to keep track of the origin of the various terms. We will keep track only of the scale variables ε, δ and not of the other dimensional factors. For example, $H^{(2)}$ itself contains the following type of terms:

$$\begin{aligned} H^{(2)} &= \varepsilon^2 \frac{1}{2} \int_{\mathbb{R}} \left(\frac{h}{\rho} \tilde{u}^2 + \frac{2h}{\rho} \tilde{u} \tilde{u}_1 + \left(\frac{h}{\rho} + \frac{h_1}{\rho_1} \right) \tilde{u}_1^2 \right) dx \\ &= \varepsilon^2 \frac{1}{2} (\rho - \rho_1) (g + \gamma \kappa) \int_{\mathbb{R}} \eta^2 dx + \varepsilon^2 \frac{1}{2} \rho_1 (g + \gamma_1^2 h_1 + \gamma_1 \kappa_1) \int_{\mathbb{R}} \eta_1^2 dx \\ &\quad + \varepsilon^2 \kappa \int_{\mathbb{R}} \tilde{u} \eta dx + \varepsilon^2 (\kappa_1 + \gamma_1 h_1) \int_{\mathbb{R}} \tilde{u}_1 \eta_1 dx \\ &\quad + \varepsilon^2 \delta^2 \frac{1}{2} \int_{\mathbb{R}} \left(-\frac{h^2}{3\rho^2} (\rho h + 3\rho_1 h_1) \tilde{u}_x^2 - \frac{h}{3\rho^2} (2\rho h^2 + 6\rho_1 h h_1 + 3\rho h_1) \tilde{u}_x \tilde{u}_{1,x} \right) dx \\ &\quad + \varepsilon^2 \delta^2 \frac{1}{2} \int_{\mathbb{R}} \left(-\frac{1}{3\rho^2 \rho_1} (\rho^2 h_1^3 + \rho \rho_1 h^3 + 3\rho \rho_1 h h_1^2 + 3\rho_1 h^2 h_1) \tilde{u}_{1,x}^2 \right) dx. \end{aligned} \tag{85}$$

Here $H^{(2)}$ is given in terms of $\tilde{u}, \eta, \tilde{u}_1, \eta_1$ which are not canonical variables. The canonical are the variables

$$z = \xi + \frac{\Gamma}{2} \partial_x^{-1} \eta, \quad z_1 = \xi_1 + \frac{\Gamma_1}{2} \partial_x^{-1} \eta_1. \tag{86}$$

It is more convenient to work however with the variables which are x -derivatives of z and z_1 :

$$p = \tilde{u} + \frac{\Gamma}{2}\eta, \quad p_1 = \tilde{u}_1 + \frac{\Gamma_1}{2}\eta_1. \quad (87)$$

In terms of these new variables, the Hamiltonian structure changes accordingly,

$$\frac{d}{dt} \begin{pmatrix} p \\ p_1 \\ \eta \\ \eta_1 \end{pmatrix} = -\varepsilon^{-2} \partial_x \begin{pmatrix} 0 & 0 & 1 & 0 \\ 0 & 0 & 0 & 1 \\ 1 & 0 & 0 & 0 \\ 0 & 1 & 0 & 0 \end{pmatrix} \begin{pmatrix} \delta H / \delta p \\ \delta H / \delta p_1 \\ \delta H / \delta \eta \\ \delta H / \delta \eta_1 \end{pmatrix}, \quad (88)$$

In terms of the old variables $\tilde{u}, \eta, \tilde{u}_1, \eta_1$

$$\frac{d}{dt} \begin{pmatrix} \tilde{u} \\ \tilde{u}_1 \\ \eta \\ \eta_1 \end{pmatrix} = -\varepsilon^{-2} \partial_x \begin{pmatrix} -\Gamma & 0 & 1 & 0 \\ 0 & -\Gamma_1 & 0 & 1 \\ 1 & 0 & 0 & 0 \\ 0 & 1 & 0 & 0 \end{pmatrix} \begin{pmatrix} \delta H / \delta \tilde{u} \\ \delta H / \delta \tilde{u}_1 \\ \delta H / \delta \eta \\ \delta H / \delta \eta_1 \end{pmatrix}, \quad (89)$$

Since $\mathcal{O}(\varepsilon^2 \delta^2) = \varepsilon^3$ then already $H^{(2)}$ in the long wave approximation produces terms of order ε^3 . For this reason we will not need terms of order $\varepsilon^3 \delta^2 \sim \varepsilon^4$ from $H^{(3)}$. Therefore, the relevant terms from $H^{(3)}$ are

$$\begin{aligned} H^{(3)} &= \varepsilon^3 \frac{1}{2} \int_{\mathbb{R}} \left(\frac{1}{\rho} \eta \tilde{u}^2 + \frac{2}{\rho} \eta \tilde{u} \tilde{u}_1 - \frac{\rho - \rho_1}{\rho \rho_1} \eta \tilde{u}_1^2 + \frac{1}{\rho_1} \eta_1 \tilde{u}_1^2 \right) dx \\ &+ \varepsilon^3 \frac{1}{2} (\rho - \rho_1) \gamma^2 \int_{\mathbb{R}} \frac{\eta^3}{3} dx + \varepsilon^3 \frac{1}{2} \rho_1 \gamma_1^2 \int_{\mathbb{R}} \frac{\eta_1^3}{3} dx \\ &+ \varepsilon^3 \frac{1}{2} \gamma \int_{\mathbb{R}} \tilde{u} \eta^2 dx + \varepsilon^3 \frac{1}{2} \gamma_1 \int_{\mathbb{R}} \tilde{u}_1 \eta_1^2 dx. \end{aligned} \quad (90)$$

The next assumption in our approximation is that the pair of canonical Hamiltonian variables p_1 and η_1 (i.e. with respect to the Hamiltonian structure in (89)) associated to the free surface are of smaller order,

$$\eta_1 \rightarrow \varepsilon \eta, \quad p_1 \rightarrow \varepsilon p_1. \quad (91)$$

Clearly $\tilde{u}_1 \rightarrow \varepsilon \tilde{u}_1$. The approximate Hamiltonian of the system is with terms of orders ε^2 and ε^3 :

$$\begin{aligned}
H_a &= \varepsilon^2 \frac{1}{2} \int_{\mathbb{R}} \left(\frac{h}{\rho} \tilde{u}^2 + \varepsilon \frac{2h}{\rho} \tilde{u} \tilde{u}_1 + (\rho - \rho_1)(g + \gamma \kappa) \eta^2 \right) dx \\
&+ \varepsilon^2 \kappa \int_{\mathbb{R}} \tilde{u} \eta dx - \varepsilon^2 \delta^2 \frac{1}{2} \int_{\mathbb{R}} \frac{h^2}{3\rho^2} (\rho h + 3\rho_1 h_1) \tilde{u}_x^2 dx \\
&+ \varepsilon^3 \frac{1}{2} \int_{\mathbb{R}} \left(\frac{1}{\rho} \eta \tilde{u}^2 \right) dx + \varepsilon^3 \frac{1}{2} (\rho - \rho_1) \gamma^2 \int_{\mathbb{R}} \frac{\eta^3}{3} dx + \varepsilon^3 \frac{1}{2} \gamma \int_{\mathbb{R}} \tilde{u} \eta^2 dx.
\end{aligned} \tag{92}$$

For the sake of simplification, let us introduce the following notations for some constants:

$$\begin{aligned}
a_1 &= (\rho - \rho_1)(g + \gamma \kappa), \\
a_2 &= \frac{h^2}{3\rho^2} (\rho h + 3\rho_1 h_1).
\end{aligned} \tag{93}$$

The equations are

$$\begin{aligned}
\eta_t &= -\varepsilon^{-2} \left(\frac{\delta H_a}{\delta \tilde{u}} \right)_x, \\
(\tilde{u} + \Gamma \eta)_t &= -\varepsilon^{-2} \left(\frac{\delta H_a}{\delta \eta} \right)_x, \\
\eta_{1,t} &= -\varepsilon^{-2} \left(\frac{\delta H_a}{\delta \tilde{u}_1} \right)_x, \\
(\tilde{u}_1 + \Gamma_1 \eta_1)_t &= -\varepsilon^{-2} \left(\frac{\delta H_a}{\delta \eta_1} \right)_x = 0.
\end{aligned} \tag{94}$$

Due to the last equation we have

$$\tilde{u}_1 = -\Gamma_1 \eta_1 \tag{95}$$

and for the other variables

$$\eta_t = - \left(\frac{h}{\rho} \tilde{u} + \kappa \eta + \varepsilon \frac{h}{\rho} \tilde{u}_1 + \delta^2 a_2 \tilde{u}_{xx} + \varepsilon \left(\frac{1}{\rho} \eta \tilde{u} + \frac{\gamma}{2} \eta^2 \right) \right)_x, \quad (96)$$

$$(\tilde{u} + \Gamma \eta)_t = - (a_1 \eta + \kappa \tilde{u})_x - \varepsilon \left(\frac{1}{2\rho} \tilde{u}^2 + \frac{(\rho - \rho_1) \gamma^2}{2} \eta^2 + \gamma \tilde{u} \eta \right)_x, \quad (97)$$

$$\eta_{1,t} = -\varepsilon \frac{h}{\rho} \tilde{u}_x. \quad (98)$$

The leading order linear equations for \tilde{u} and η are

$$\begin{aligned} \eta_t &= - \left(\frac{h}{\rho} \tilde{u}_x + \kappa \eta_x \right), \\ (\tilde{u} + \Gamma \eta)_t &= - (a_1 \eta_x + \kappa \tilde{u}_x). \end{aligned} \quad (99)$$

The wavespeed c of the solutions, proportional to $e^{ik(x-ct)}$ satisfies the quadratic equation

$$(c - \kappa)^2 + \frac{h\Gamma}{\rho}(c - \kappa) + \frac{h}{\rho}(\kappa\Gamma - a_1) = 0. \quad (100)$$

Introducing $c_0(k) = c(k) - \kappa$ and noting that

$$a_1 - \kappa\Gamma = (\rho - \rho_1)(g - 2\omega\kappa) \approx g(\rho - \rho_1)$$

we write the equation for $c_0(k)$ in the form

$$c_0^2 + \frac{h\Gamma}{\rho}c_0 - \frac{\rho - \rho_1}{\rho}gh = 0. \quad (101)$$

The solution is

$$c_0 = \frac{1}{2} \left(-\frac{h\Gamma}{\rho} \pm \sqrt{\left(\frac{h\Gamma}{\rho}\right)^2 + 4\frac{\rho - \rho_1}{\rho}gh} \right). \quad (102)$$

There are right ($c_0 > 0$) and left ($c_0 < 0$) running waves. We notice that in this approximation $c(k)$ is k -independent, i.e. constant.

In the zero vorticity case

$$c \rightarrow c' = \kappa \pm \sqrt{\frac{\rho - \rho_1}{\rho}} gh. \quad (103)$$

Let us introduce also $c'_0 = c' - \kappa$. From (99) it also follows that in the leading order $\tilde{u} = (\rho c_0 \eta)/h$, i.e.

$$\tilde{u} = \frac{\rho c_0}{h} \eta + \varepsilon q + \mathcal{O}(\varepsilon^2) \quad (104)$$

for some yet unknown quantity q . Nevertheless from (98) and (104) we determine

$$\eta_1 = \varepsilon \frac{c_0}{c} \eta + \mathcal{O}(\varepsilon^2), \quad (105)$$

$$\tilde{u}_1 = -\Gamma_1 \eta_1 = -\varepsilon \frac{\Gamma_1 c_0}{c} \eta + \mathcal{O}(\varepsilon^2). \quad (106)$$

Note that $\mathcal{O}(\tilde{u}_1) = \varepsilon$, thus $\mathcal{O}(\varepsilon \tilde{u}_1) = \varepsilon^2$ and such terms will be neglected in (96) and (97). The most general form for q is

$$\varepsilon q = \varepsilon b_1 \eta^2 + \delta^2 b_2 \eta_{xx} + \mathcal{O}(\varepsilon^2) \quad (107)$$

for some constants b_1, b_2 . Now we are in a position to express everything in (96) and (97) only via the variable η . This way we obtain two evolutionary equations for η which should coincide up to $\mathcal{O}(\varepsilon)$. The equality of their coefficients allows us to find

$$b_1 = \frac{\rho c_0 (\gamma - \frac{c_0}{h}) - 2h\omega(\rho - \rho_1)\gamma - 2\Gamma c_0}{2h \left(2c_0 + \frac{h\Gamma}{\rho} \right)}, \quad (108)$$

$$b_2 = \frac{-a_2 \rho^2 c_0 (\rho c_0 + h\Gamma)}{h^2 (2\rho c_0 + h\Gamma)}. \quad (109)$$

η satisfies the KdV equation [39], (see also [10, 35])

$$\eta_t + c\eta_x + \delta^2 \left(\frac{h}{\rho} b_2 + \frac{a_2 \rho c_0}{h} \right) \eta_{xxx} + \varepsilon \left(\frac{h}{\rho} b_1 + \frac{c_0}{h} + \frac{\gamma}{2} \right) 2\eta\eta_x = 0 \quad (110)$$

which, in the limit when all vorticities are zero becomes

$$\eta_t + c'\eta_x + \delta^2 \frac{\rho c_0' h}{6\rho} (\rho h + 3\rho_1 h_1) \eta_{xxx} + \varepsilon \frac{3c_0'}{2h} \eta \eta_x = 0. \quad (111)$$

From η one can recover \tilde{u} and as well as η_1 and \tilde{u}_1 :

$$\tilde{u} = \frac{\rho c_0}{h} \eta + \varepsilon b_2 \eta^2 + \delta^2 b_3 \eta_{xx} + \mathcal{O}(\varepsilon^2), \quad (112)$$

$$\eta_1 = \varepsilon \frac{c_0}{c} \eta + \mathcal{O}(\varepsilon^2), \quad (113)$$

$$\tilde{u}_1 = -\varepsilon \frac{\Gamma_1 c_0}{c} \eta + \mathcal{O}(\varepsilon^2). \quad (114)$$

We notice that in leading order

$$\frac{\eta_1}{\eta} = \varepsilon \frac{c_0}{c} \sim \frac{c_0}{c_0 + \kappa}$$

and therefore both positive and negative relative polarities for η and η_1 are possible. The KdV approximation for an internal wave coupled to a free surface for a different configuration of the currents is derived in [15].

The KdV equation represents a balance between a nonlinearity term $\varepsilon \eta \eta_x$, and dispersion term $\delta^2 \eta_{xxx}$. In the above considerations ε and δ^2 , are of the same order and as a result we can have the stable soliton solutions of the KdV equation. However, there are various geophysical scales and other situations are possible, including $\delta \sim \varepsilon^2$. In such case $\delta^2 \sim \varepsilon^4 \ll \varepsilon$ and instead of the KdV equation the relevant model is the dispersionless Burgers equation ($\partial_\tau = \partial_t + c \partial_x$)

$$\eta_\tau + \varepsilon \left(\frac{h}{\rho} b_1 + \frac{c_0}{h} + \frac{\gamma}{2} \right) 2\eta \eta_x = 0. \quad (115)$$

It is well known that the solutions of this equation always form a vertical slope and break. Such wave-breaking phenomenon is well known for internal waves in the ocean. This is a mechanism that causes mixing in the deep ocean, [40].

There are other integrable systems which provide an approximation of the equations in the Boussinesq regime, such as the Kaup-Boussinesq system investigated firstly by D.J. Kaup [38], see also [33]. Two-component

integrable systems, that can match the model equations up to order δ^2 , are the 2-component Camassa-Holm system and the Zakharov-Ito system [14, 32, 31, 25, 23, 24].

9. Discussion and conclusions

We consider a two-media system of liquids with different densities, free surface and a free internal surface separating the liquids. The bottom of the system is considered horizontal and flat. We studied the surface and internal waves driven by gravity and Coriolis forces and interacting with a current. The underlying current is in the form of a shear flow with a specific velocity profile. The current has constant vorticity at the horizontal strips where the surface and internal waves are localised. The model is aimed at geophysical applications, where a typical configuration is the one of a thin shallow layer of warm and less dense water over a much deeper layer of cold denser water. The governing equations are written in a canonical Hamiltonian form, which gives rise to a systematic approach for possible approximations. In particular, small amplitude and long-wave regimes are studied. There are various geophysical scales, allowing for smooth solitons at the KdV regime as well as breaking waves in the very large wavelength regime, when the equations are asymptotically equivalent to the dispersionless Burgers equation. In the case of a free surface, even in the case of very small amplitudes, the internal wave is usually coupled to the surface wave. This has an impact on the propagation speeds and is observed in other related models [16, 15]. Other asymptotic regimes, e.g. related to the Nonlinear Schrödinger equation [51] remain to be studied. The stability of the waves interacting with currents is another important issue that needs to be addressed in the future. Related recent works in this connection are [13, 26, 30].

Acknowledgements

The author acknowledges Seed funding grant support from Dublin Institute of Technology for a project in association with the Environmental Sustainability and Health Institute, Dublin. The author is grateful to Prof. Adrian Constantin for many valuable discussions and to an anonymous referee for many important suggestions.

References

- [1] T. Benjamin, T. Bridges, *Reappraisal of the Kelvin-Helmholtz problem. Part 1. Hamiltonian structure*. J. Fluid Mech. **333** (1997), 301–325.
- [2] T. Benjamin, T. Bridges, *Reappraisal of the Kelvin-Helmholtz problem. Part 2. Interaction of the Kelvin-Helmholtz, superharmonic and Benjamin-Feir instabilities*. J. Fluid Mech. **333** (1997), 327–373.
- [3] T. Benjamin, P. Olver, *Hamiltonian structure, symmetries and conservation laws for water waves*. J. Fluid Mech. **125** (1982), 137–185.
- [4] A. Compelli, *Hamiltonian formulation of 2 bounded immiscible media with constant non-zero vorticities and a common interface*. Wave Motion **54** (2015), 115–124.
- [5] A. Compelli, *Hamiltonian approach to the modeling of internal geophysical waves with vorticity*. Monatsh. Math. **179(4)** (2016), 509–521.
- [6] A. Compelli, R. Ivanov, *On the dynamics of internal waves interacting with the Equatorial Undercurrent*. J. Nonlinear Math. Phys. **22** (2015), 531–539; arXiv:1510.04096 [math-ph]
- [7] A. Compelli, R. Ivanov, *Hamiltonian approach to internal wave-current interactions in a two-media fluid with a rigid lid*. Pliska Stud. Math. Bulgar. **25** (2015) 7–18, arXiv:1607.01358 [physics.flu-dyn]
- [8] A. Compelli, R. I. Ivanov, *The Dynamics of Flat Surface Internal Geophysical Waves with Currents*. Journal of Mathematical Fluid Mechanics (2016) DOI: 10.1007/s00021-016-0283-4; arXiv:1611.06581 [physics.flu-dyn]
- [9] A. Constantin, *On the deep water wave motion*. J. Phys. A **34** (2001), 1405–1417.
- [10] A. Constantin, *Nonlinear water waves with applications to wave-current interactions and tsunamis*. CBMS-NSF Regional Conference Series in Applied Mathematics, **81**, SIAM, Philadelphia, PA, 2011.
- [11] A. Constantin, J. Escher, *Symmetry of steady periodic surface water waves with vorticity*. J. Fluid Mech. **498** (2004), 171–181.

- [12] A. Constantin, J. Escher, *Analyticity of periodic traveling free surface water waves with vorticity*. Ann. of Math. **173** (2011), 559–568.
- [13] A. Constantin, P. Germain, *Instability of some equatorially trapped waves*. J. Geophys. Res. Oceans **118** (2013), 2802–2810.
- [14] A. Constantin, R. Ivanov, *On an integrable two-component Camassa-Holm shallow water system*, Physics Letters A. **372** (2008), 7129–7132; arXiv:0806.0868 [nlin.SI]
- [15] A. Constantin, R. Ivanov, *A Hamiltonian approach to wave-current interactions in two-layer fluids*. Phys. Fluids **27** (2015), 086603.
- [16] A. Constantin, R. Ivanov, C.-I. Martin, *Hamiltonian formulation for wave-current interactions in stratified rotational flows*. Arch. Rational Mech. Anal. **221** (2016), 1417–1447, DOI: 10.1007/s00205-016-0990-2.
- [17] A. Constantin, R. Ivanov, E. Prodanov, *Nearly-Hamiltonian structure for water waves with constant vorticity*. J. Math. Fluid Mech. **9** (2007), 1–14; arXiv:math-ph/0610014
- [18] A. Constantin, R.S. Johnson, *The dynamics of waves interacting with the Equatorial Undercurrent*. Geophysical & Astrophysical Fluid Dynamics **109** (2015) 311–358.
- [19] A. Constantin, R. S. Johnson, *An exact, steady, purely azimuthal equatorial flow with a free surface*. J. Phys. Oceanogr. **46** (2016), 1935–1945.
- [20] C. Cotter, D. Holm, J. Percival, *The square root depth wave equations*. Proc. R. Soc. A **466** (2010) 3621–3633, doi:10.1098/rspa.2010.0124
- [21] W. Craig, P. Guyenne, H. Kalisch, *Hamiltonian long wave expansions for free surfaces and interfaces*. Comm. Pure Appl. Math. **24** (2005), 1587–1641.
- [22] W. Craig, P. Guyenne, C. Sulem, *Coupling between internal and surface waves*, Nat. Hazards **57** (2011) 617–642.
- [23] J. Escher, D. Henry, B. Kolev, T. Lyons, *Two-component equations modelling water waves with constant vorticity*. Annali di Matematica Pura ed Applicata **195** (2016), 249–271.

- [24] J. Escher, O. Lechtenfeld and Z. Yin, *Well-posedness and blow-up phenomena for the 2-component Camassa-Holm equation*. Discrete Contin. Dyn. Syst. **19** (2007), 495–513.
- [25] L.Fan, H. Gao and Y.Liu, *On the rotation-two-component Camassa-Holm system modelling the equatorial water waves*. Adv. Math. **291** (2016), 59–89.
- [26] F. Genoud, and D. Henry, *Instability of equatorial water waves with an underlying current*. J.Math. Fluid Mech. **16** (2014), 661–667.
- [27] D. Henry, *An exact solution for Equatorial geophysical water waves with an underlying current*. Eur. J. Mech. B Fluids **38** (2013), 190–195.
- [28] D. Henry, *Internal Equatorial water waves in the f -plane*. J. Nonlinear Math. Phys. **22** (2015), 499–506.
- [29] D. Henry, *Exact equatorial water waves in the f -plane*. Non. Anal. Real World Appl. **28** (2016), 284–289.
- [30] D. Henry, H.-C. Hsu, *Instability of equatorial water waves in the f -plane*. Discrete Contin. Dyn. Syst. **35** (2015), 906–916.
- [31] D. D. Holm and R. I. Ivanov, *Two-component CH system: Inverse scattering, peakons and geometry*. Inverse Problems **27** (2011), 045013, arXiv:1009.5374v1 [nlin.SI].
- [32] R. Ivanov, *Two component integrable systems modelling shallow water waves: the constant vorticity case*. Wave Motion **46** (2009), 389–396; arXiv:0906.0780 [nlin.SI]
- [33] R. Ivanov and T. Lyons, *Integrable models for shallow water with energy dependent spectral problems*. Journal of Nonlinear Mathematical Physics **19**, Suppl. 1 (2012), 1240008 (17 pages); arXiv:1211.5567 [nlin.SI]
- [34] T. Izumo, *The equatorial current, meridional overturning circulation, and their roles in mass and heat exchanges during the El Niño events in the tropical Pacific ocean*, Ocean Dynamics 55 (2005), 110–123.
- [35] R.S. Johnson, *Camassa-Holm, Korteweg-de Vries and related models for water waves*. J. Fluid. Mech. **457** (2002), 63–82.

- [36] G. C. Johnson, M. J. McPhaden and E. Firing, *Equatorial Pacific Ocean horizontal velocity, divergence, and upwelling*. J. Phys. Oceanogr. **31** (2001), 839–849.
- [37] I. G. Jonsson, *Wave-current interactions*. In: The Sea (Wiley, New York, 1990) pp 65–120.
- [38] D.J. Kaup, *A higher-order water-wave equation and the method for solving it*. Progr. Theoret. Phys. **54** (1975), 396–408.
- [39] D. Korteweg, G. de Vries, *On the change of form of long waves advancing in a rectangular canal, and on a new type of long stationary waves*. Philosophical Magazine **39** (1895), 422–443, reprint: Philosophical Magazine **91** (2011), 1007–1028. DOI: 10.1080/14786435.2010.547337
- [40] K. Lamb, *Internal Wave Breaking and Dissipation Mechanisms on the Continental Slope/Shelf*, Annual Review of Fluid Mechanics **46** (2014), 231–254.
- [41] C.-I. Martin, *Dispersion relations for gravity water flows with two rotational layers*. Eur. J. of Mech. B/Fluids **50** (2015), 9–18.
- [42] C.-I. Martin, B.-V. Matioc, *Steady Periodic Water Waves with Unbounded Vorticity: Equivalent Formulations and Existence Results*. J. Nonlinear Science **24** (2014), 633–659.
- [43] S. A. Maslowe, *Critical layers in shear flows*. Annu. Rev. Fluid Mech. **18** (1986), 405–432.
- [44] A. V. Matioc, *An exact solution for geophysical equatorial edge waves over a sloping beach*. J. Phys. A **45** (2012), 365501.
- [45] A. V. Matioc, *Exact geophysical waves in stratified fluids*. Appl. Anal. **92** (2013), 2254–2261.
- [46] D. Milder, *A note regarding “On Hamilton’s principle for water waves”*. J. Fluid Mech. **83** (1977), 159–161.
- [47] J. Miles, *Hamiltonian formulations for surface waves*. Appl. Sci. Res. **37** (1981), 103–110.

- [48] J. Miles, *On Hamilton's principle for water waves*. J. Fluid Mech. **83** (1977), 153–158.
- [49] D. H. Peregrine, *Interaction of water waves and currents*. Adv. Appl. Mech. **16** (1976), 9–117.
- [50] A. Teles da Silva, D. Peregrine, *Steep, steady surface waves on water of finite depth with constant vorticity*. J. Fluid Mech. **195** (1988), 281–302.
- [51] R. Thomas, C. Kharif and M. Manna, *A nonlinear Schrödinger equation for water waves on finite depth with constant vorticity*. Physics of fluids, **24** (2012), 127102.
- [52] G. P. Thomas and G. Klopman, *Wave-current interactions in the near-shore region*. In: *Gravity waves in water of finite depth*, edited by J. N. Hunt (computational Mechanics, Southampton, 1997) 215–319.
- [53] E. Wahlén, *A Hamiltonian formulation of water waves with constant vorticity*. Lett. Math. Phys. **79** (2007), 303–315.
- [54] V. Zakharov, *Stability of periodic waves of finite amplitude on the surface of a deep fluid*. Zh. Prikl. Mekh. Tekh. Fiz. **9** (1968), 86–94 (in Russian); J. Appl. Mech. Tech. Phys. **9** (1968), 190–194 (English translation).